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6 A European endeavor for optimizing
7 Zonal Isolation, Drilling and Exploitation of
8 EGS projects
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12 **REPORT AND RESULTS OF ETHZ-INVOLVED SUB-**
13 **TASKS (2.1, 3.1, 3.2, 3.4, 4.2, 4.3)**
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68 fractures (NF) are shown as well. Note that the dry reopening tests were not successful in all intervals.
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72 kink in the record, showing a decrease in the system stiffness, is taken as the fracture reopening
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81 occurring in borehole SB2.2 is marked, the other fractures are from SB3.1. The solid line gives the
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90 borehole is indicated, as well as trends in azimuth and width for groups of breakouts identified,
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108 experiment. (Kong et al., (2021)) 17

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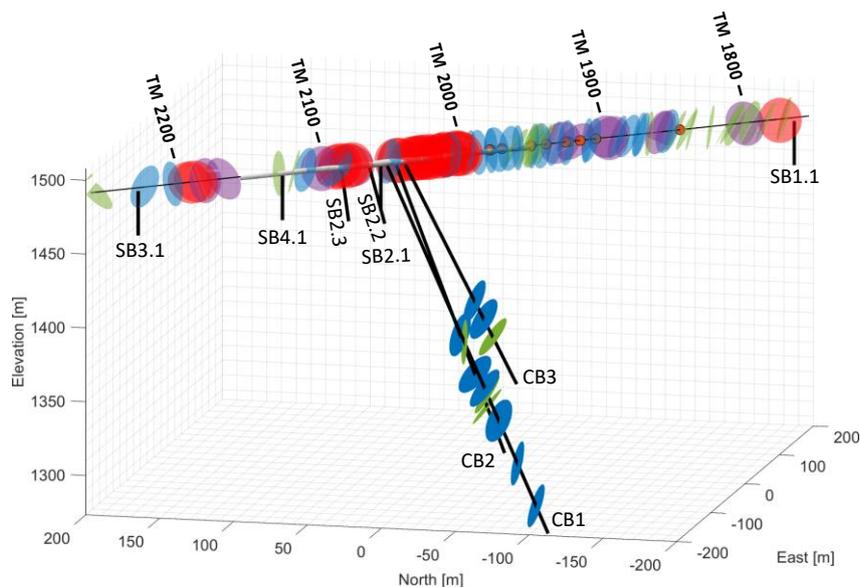
137 1. Introduction

138 This report presents the borehole characterization work at the Bedretto Underground Laboratory for
 139 Geosciences and Geoenergies (BULGG) and the flow-through experiments at the GREAT Visualization
 140 Lab of the Geothermal Energy and Geofluids Group (GEG) at ETH Zurich (ETHZ). ETHZ was involved in
 141 the borehole characterization in WP 2 (Demonstration of wellbore segmenting technologies), seismic
 142 risk and monitoring, data analysis and modelling, and demonstration of simulation with different zone
 143 isolation systems in WP 3 (Demonstration of reservoir simulation technologies and lifting
 144 optimization), and sampling and analysis of effluents from plant in WP 4. However, the involvement
 145 of ETH Zurich in WP 4 was minimized due to the suspending of collaboration between ETHZ, GES, ESG
 146 in WP 4. This deviation has been approved by Swiss Bundesamt für Energie (BFE).

147 2. Methodology

148 2.1 Field campaign at BULGG

149 For the field campaign at the BULGG, ETHZ used the Acoustic Televiwer (ATV) and Optical Televiwer
 150 (OTV) to log the boreholes (Figure 1). ETHZ analysed the stress state at the BULGG by using mini-frac
 151 tests, and dry (packer) reopening tests. The fracture closure pressure was analysed with several
 152 techniques, including G-function, square root of time, bilinear pressure-decay, and jacking pressure.
 153 ETHZ performed the step-rate injection tests to determine the fracture and matrix injectivity in the
 154 BUL. ETHZ conducted analyses of the dilatometer test to obtain the modulus of deformation and
 155 modulus of elasticity of the in-situ rock mass in the BUL. Using the ATV and caliper logs, ETH Zurich
 156 conducted borehole breakout analysis. ETHZ also contributed to the design and construction of the
 157 sensors (Streamer 2.0) of the seismic monitoring network in the BULGG.

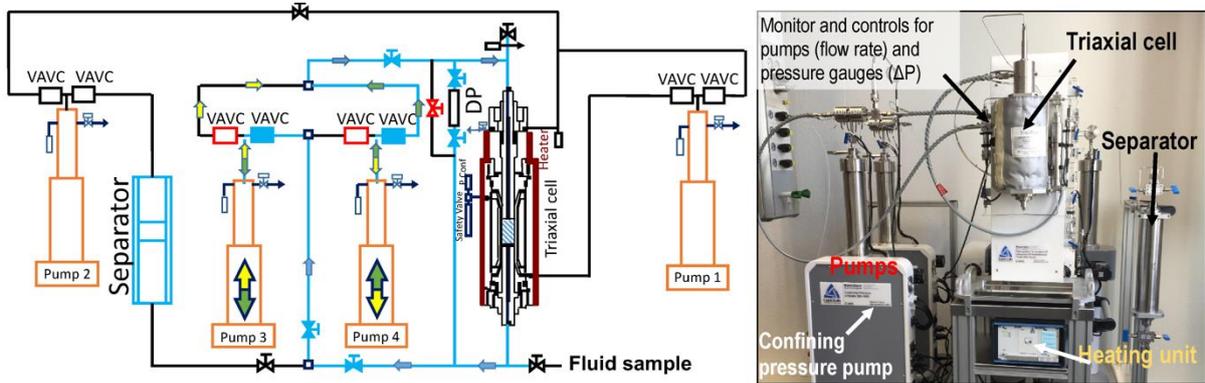


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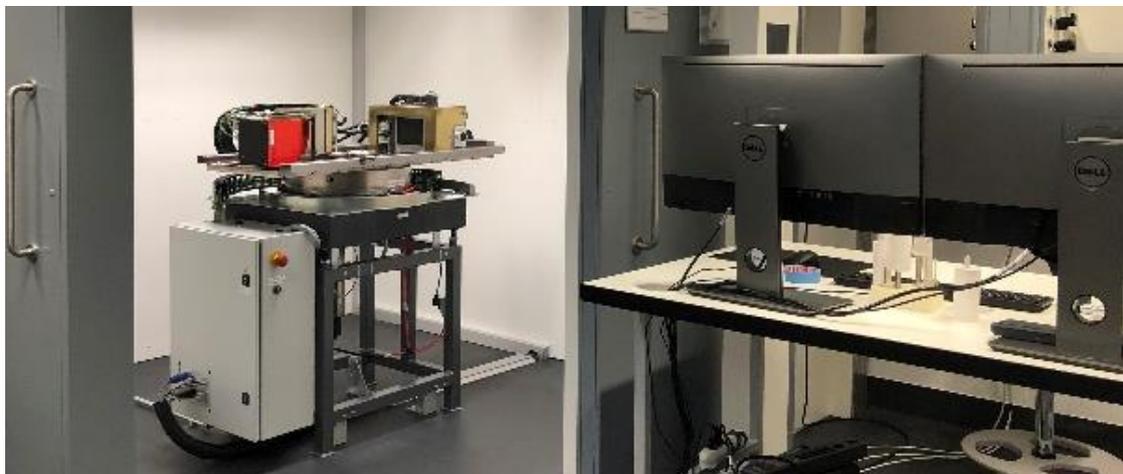
159 *Figure 1. All ETHZ works in WP 2 used boreholes in the BULGG. Analyses of borehole fractures and*
 160 *stress were done with Boreholes SB1.1, SB2.2 and SB3.1, and breakout analysis (additional work) was*
 161 *done with Boreholes CB1 and CB2.*

162 **2.2 Laboratory activities at the GEG Great Visualization Lab**

163 At the GEG Group lab, ETHZ has conducted a series of flow-through experiments. These experiments
 164 were conducted using the novel (hydrothermal) flow-through (reactor) system (Figure 2), which
 165 employs syringe pumps (up to 100 MPa and a volume of 300 ml) to provide constant fluid flow rates
 166 (from 10^{-4} ml/min to 73 ml/min) through core samples located in a dual-end- access triaxial cell.
 167 Syringe pumps also maintain outlet pore pressure (i.e., back pressure) and core (pore-fluid)
 168 confinement pressure. At the same time, electric band heaters precisely control temperature (up to
 169 200°C) of this triaxial cell. Pressure transducers, installed at the inlet and outlet of the flow-through
 170 cell, continuously monitor the differential pressure. To identify the mineral compositions of the rock
 171 samples, ETHZ performed Energy Dispersive X-Ray Analysis (EDX) and Scanning Electron Microscope
 172 (SEM) analyses. To image the rock pore/fracture geometry, ETHZ performed X-ray Computed
 173 Tomography (XRCT) scanning of rock samples (Figure 3). To obtain water chemistry (major ion
 174 concentrations) of the effluents, ETH Zurich conducted Inductively Coupled Plasma-mass-
 175 spectrometry (ICPMS) measurements.



176
 177 *Figure 2. (Right) Reactive flow-through system (reactor) for investigations of chemical and physical*
 178 *processes that affect pore-scale geometries in (fractured) rocks over a wide range of temperatures and*
 179 *(pore-fluid) pressures. (Left) Schematic of our novel flow flow-through system that allows automatic*
 180 *switching between injection pump and back- pressure pump (Pumps 3 and 4) without changing the*
 181 *flow direction.*

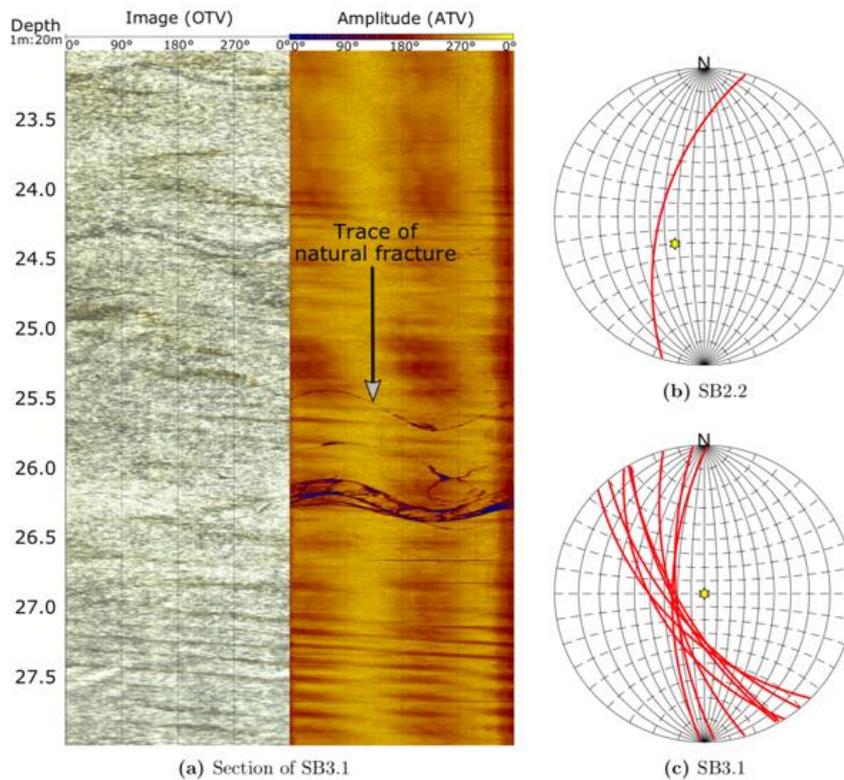


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 183 *Figure 3. The X-ray Computed Tomography scanner used for rock sample scanning at the GEG Group.*

184 **3. Results**

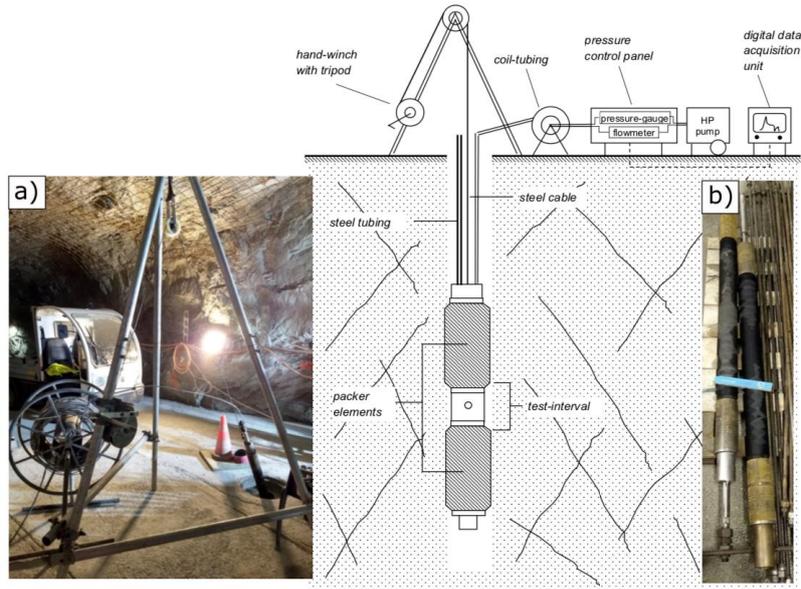
185 **3.1 Borehole analyses**

186 ETHZ has completed the proposed work in borehole characterization of WP 2: Zonal isolation. Because
 187 of the suspending of collaboration between ETH, GES, and ESG in WP 4 (see explanation on WP 4),
 188 ETH Zurich used part of the personnel hours in WP4 to conduct additional work in WP 2. This deviation
 189 has been approved by the Swiss Bundesamt für Energie (BFE). In summary, ETHZ has worked on
 190 fracture analysis, stress analysis, injectivity test, measurements of deformation modulus and elasticity
 191 modulus, and breakout analysis in WP 2. Here, the work of breakout analysis belongs to the additional
 192 work. All works in WP 2 used boreholes in the BULGG. Analyses of borehole fractures and stress were
 193 done with Boreholes SB1.1, SB2.2 and SB3.1, and breakout analysis (additional work) was done with
 194 Boreholes CB1 and CB2 (Figure 1).



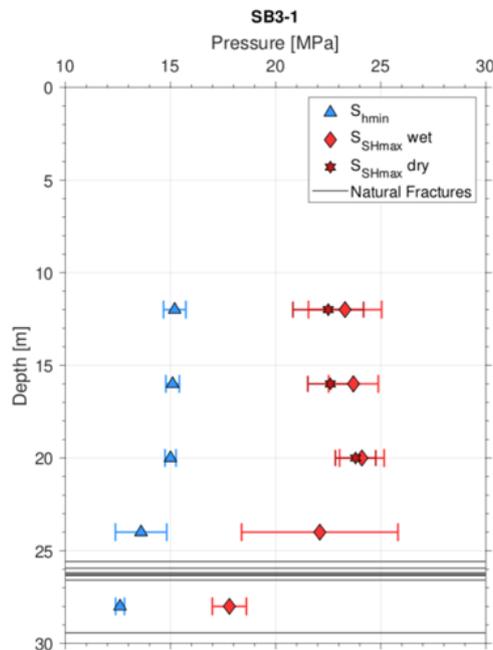
195
 196 *Figure 4. (a) Borehole logging data, ATV and OTV, showing a section with several natural fractures at*
 197 *a depth between 25.5 to 26.5 m in SB.3.1. (b) + (c) Lower hemisphere stereonet with fracture planes of*
 198 *logged natural fractures. The yellow star symbolises the borehole axis. (Bröker, 2019)*

199 Natural and/or induced fractures were identified using the logging images from Boreholes SB2.2 and
 200 SB3.1 with the Acoustic Televiwer (ATV) and Optical Televiwer (OTV) (Bröker, 2019). Borehole SB2.2
 201 shows just a single fracture at a depth of 15 m. The fracture dips into a westward direction. More
 202 fractures are visible in Borehole SB3.1, but they are limited to a depth of 25 to 27 m and a single
 203 fracture at a depth of 29 m. The dip directions vary between W and SW, with strikes between 135 ° to
 204 196 °. Dip angles of fractures in Borehole SB3.1 lie between 54 ° and 71 ° (Figure 4).



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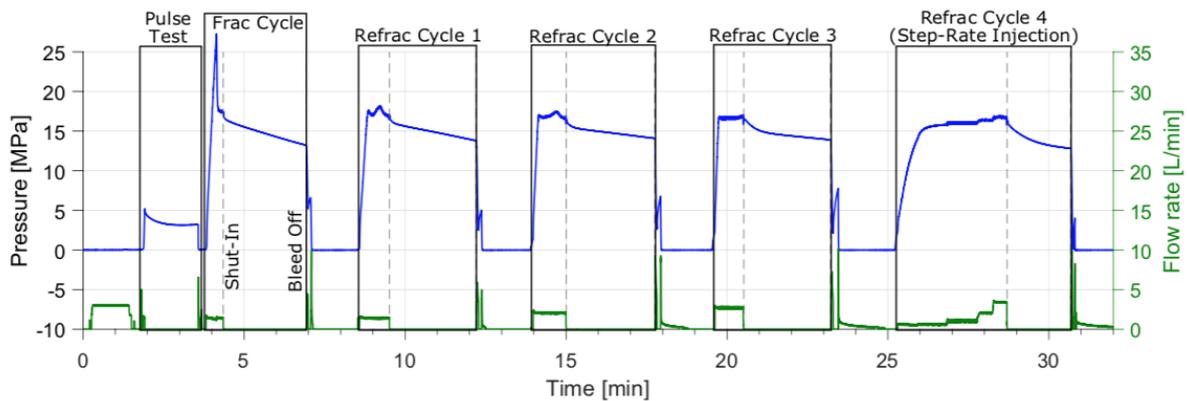
206 Figure 5. Schematic drawing of the mini-frac equipment and set up (by G. Klee, MeSy-Solexperts
 207 GmbH). Inserted are two pictures from the test site: (a) Tripod, hand-driven winch and coil-tubing;
 208 packer can be seen inside the borehole; (b) Straddle packer elements (left) and steel tubing (right).
 209 (Bröker, 2019)



210

211 Figure 6. The computed minimum (S_{hmin}) and maximum (S_{Hmax}) horizontal stresses at five depth
 212 intervals in the vertical borehole SB3.1. S_{hmin} is obtained from the fracture closure pressure picked by
 213 the compliance method, errors are given by the standard deviation per interval. S_{Hmax} is calculated from
 214 breakdown, reopening, pore pressure and S_{hmin} itself. Grey lines show where natural fractures intersect
 215 the borehole. (Bröker, 2019)

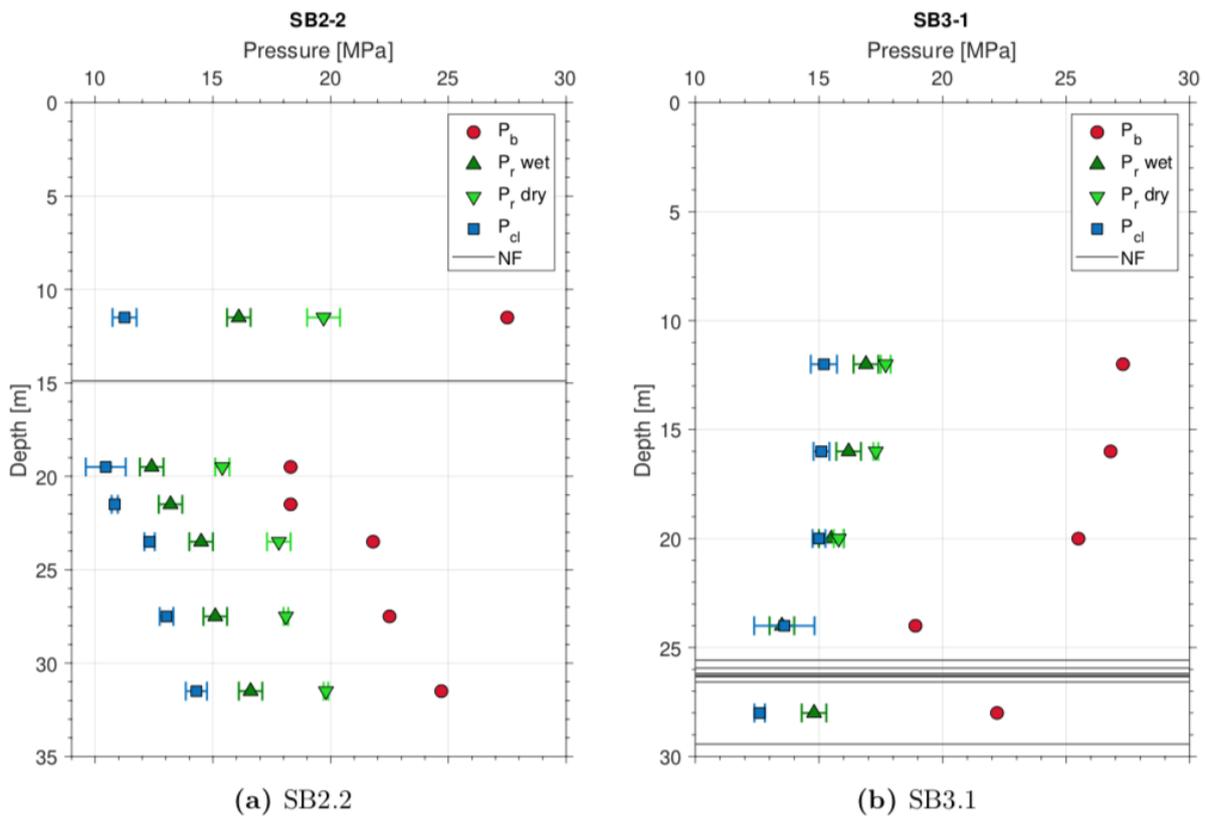
216 The stress state at the BUL was estimated with mini-frac tests and dry (packer) reopening tests (Figure
 217 5). The stresses were measured primarily based on the classic hydraulic fracturing technique, called a
 218 mini-frac test. In total, the mini-frac tests were performed in three SB boreholes (i.e., SB1.1, SB2.1,
 219 and SB3.1) at five depth intervals. If the instantaneous shut-in pressures in the mini-frac tests are used
 220 to derive the magnitude of S_{hmin} , the values of S_{hmin} falls within the range of 13-16 MPa, which is close
 221 to the stress magnitude bounded by the friction equilibrium ($S_{hmin} \sim 15.3$ MPa) with a granite density
 222 of 2.7 g/cm^3 and a frictional coefficient (μ) of 0.6 (Figure 6). In each borehole, we performed at least
 223 one over-night (typically 12-15 hours) shut-in test to measure the formation pore pressure. Several
 224 such shut-in tests indicate that pore fluid pressure (P_p) is between 3 and 6 MPa at the depths of the
 225 stress measurements. The estimation of S_{Hmax} can be derived from the difference between the
 226 breakdown pressure ($P_{breakdown}$) and the re-opening pressure ($P_{re-opening}$), e.g., $S_{Hmax} = 3S_{hmin} - P_{re-}$
 227 $opening - P_p$. Based on such a relationship, the estimation is that S_{Hmax} is about $0.8 \sim 1$ of S_v , indicating a
 228 normal faulting and/or strike-slip stress environment. If the slip tendencies of both boreholes are
 229 summarized using a Mohr circle diagram, compared to the N-S oriented fractures, the NW-SE oriented
 230 fractures exhibit higher slip tendency, which can be explained by their orientation relative to the input
 231 direction of the minimum horizontal stress.



232
 233 *Figure 7. Injection protocol of the mini-frac tests, consisting of an initial pulse test, the*
 234 *fracking/breakdown cycle and four reopening cycles, where the last one is a step-rate injection test.*
 235 *The blue curve shows the pressure inside the interval and the green one the injection flow rate. Black*
 236 *boxes mark the time windows, where fluid is injected into the interval and the pressure transient*
 237 *measured after pumping stopped (dashed line). The shown data are from SB3.1 at 12 m depth. (Bröker,*
 238 *2019)*

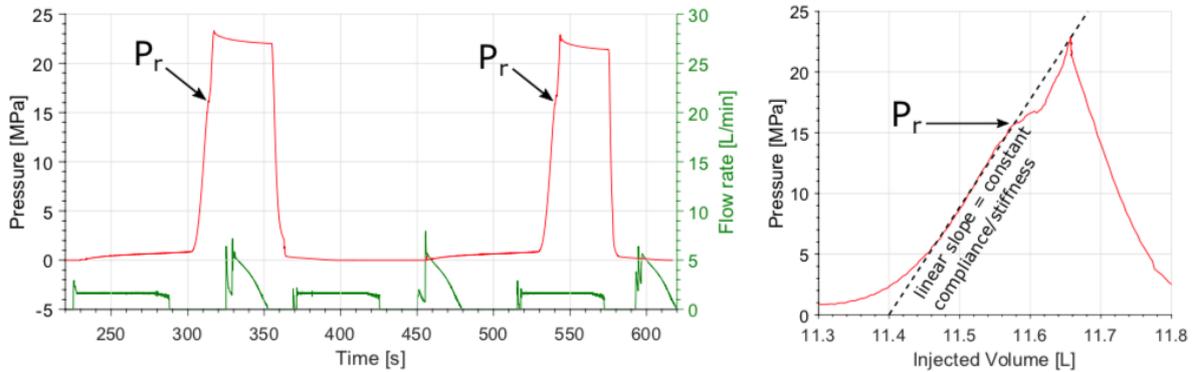
239 To increase the accuracy of fracture reopening pressure measurements and estimate the stiffness of
 240 the rock mass, dry (packer) reopening tests were conducted. The technique is based on the so-called
 241 sleeve fracturing method. Inside a borehole, a packer is inflated and pressure onto the borehole wall
 242 exerted via its membrane. By increasing the pressure further, formation breakdown can be measured,
 243 similar to hydraulic fracturing (Figure 8). Consecutive cycles can be used to assess the fracture
 244 reopening pressure (Figure 9). The tests can be used to determine the minimum and maximum
 245 horizontal stresses in laboratory tests and vertical wells. The fracture closure pressure was analysed
 246 with several techniques, including G-function, square root of time, bilinear pressure-decay, and
 247 jacking pressure. In most cases, the applied techniques give consistent results within an uncertainty
 248 range of 1 to 2 MPa. The fracture compliance method was used to identify the point, where the
 249 stiffness of the fracture increases, related to the beginning of its closure. Increasing fracture stiffness

250 was correlated to a linear or bilinear flow regime on log-log scale plots of the pressure derivative.
 251 Formation breakdown, fracture reopening and instantaneous shut-in pressures were picked and
 252 compared as well. The results show that intra- and inter-borehole variations and the proximity to
 253 naturally fractured regions seem to influence the data quality and pressure values. Natural fractures,
 254 identified from on the ATV and OTV borehole logs, are expected to be reactivated in the pressure
 255 range reached during the mini-frac tests. As the induced tensile fracture propagates further away from
 256 the borehole with every injection cycle, it becomes more and more likely that it intersects pre-existing
 257 fractures. This can be seen as multiple closure signature on several of the used diagnostic plots, where
 258 it is beneficial to have extended observation times (≥ 1 h) to fully characterise the different closure
 259 behaviours.



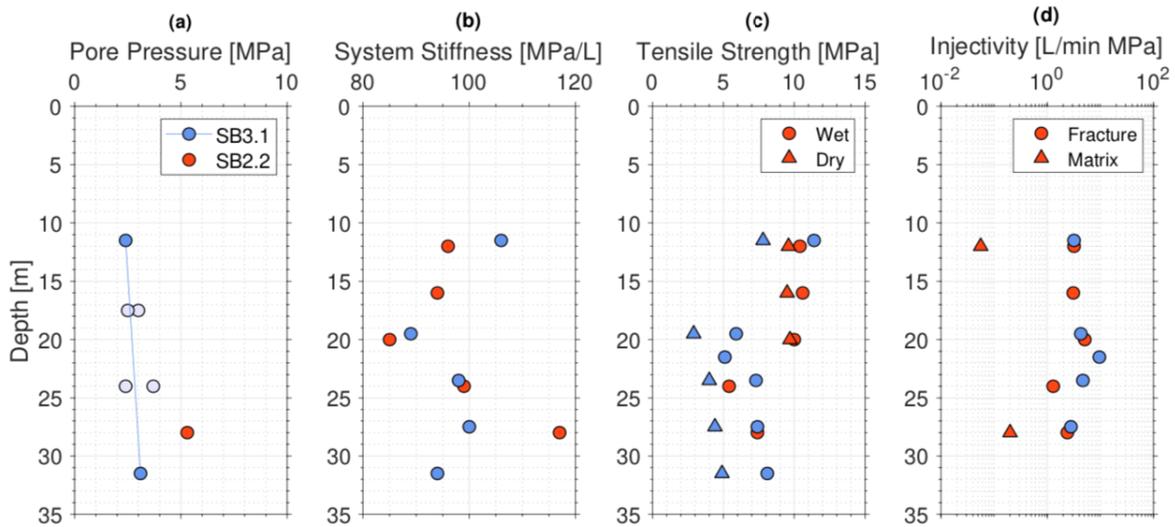
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261 *Figure 8. Results for the formation breakdown, fracture reopening and closure pressure along the two*
 262 *boreholes. The estimated errors are given by the standard deviation or are assumed. Traces of natural*
 263 *fractures (NF) are shown as well. Note that the dry reopening tests were not successful in all intervals.*
 264 *P_{cl} is the fracture closure pressure. P_b is the formation breakdown pressure. P_r is the fracture reopening*
 265 *pressure. (Bröker, 2019)*



266

267 *Figure 9. Dry reopening test in SB3.1 at 20m depth. In both plots, the deviation from a linear trend, or*
 268 *kink in the record, showing a decrease in the system stiffness, is taken as the fracture reopening*
 269 *pressure. (Left) Plot of pressure vs. time for the second and third dry reopening cycles. (Right) Plot of*
 270 *pressure vs. injected volume for the third cycle. (Bröker, 2019)*



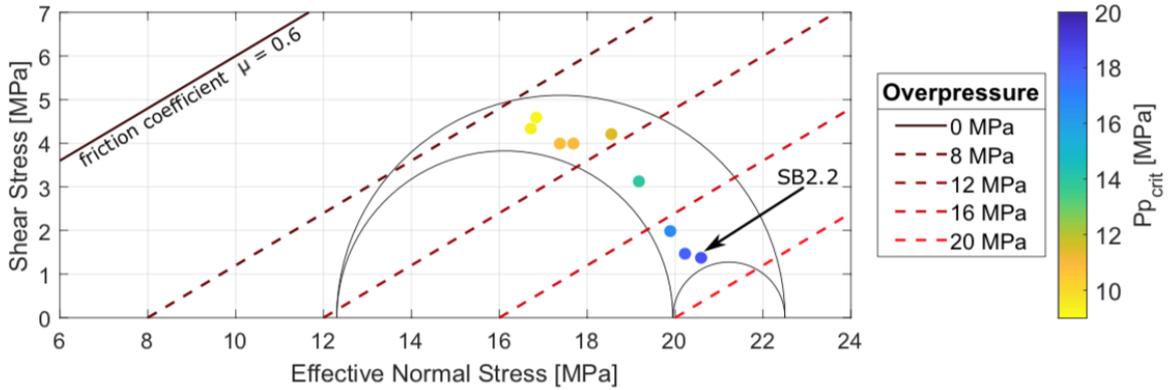
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272 *Figure 10. (a) Pore pressure evaluated from extended overnight shut-ins. (b) System stiffness calculated*
 273 *from dry reopening tests. (c) Hydraulic tensile strength calculated from formation breakdown and wet*
 274 *or dry reopening pressures. (d) Injectivity of fracture and matrix estimated from step-rate injection*
 275 *tests. (Bröker, 2019)*

276 The fracture and matrix injectivity were estimated using the step-rate injection tests (Figure 10).
 277 Matrix injectivity could just be determined in two intervals in the SB3.1 borehole, due to insufficient
 278 data points that indicate pressure versus flow rate. Both injectivity values are on the same order of
 279 magnitude with 0.06 and 0.20 L/(min MPa). The fracture injectivity was estimated for all step-rate
 280 tests, the average value in SB2.2 borehole is 3.7 L/(min MPa) and in SB3.1 borehole is 3.0 L/(min MPa).

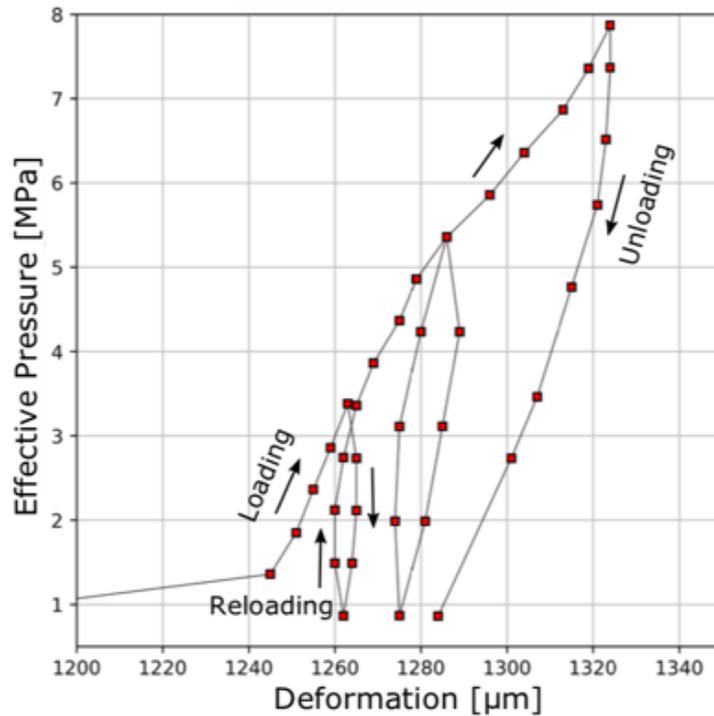
281 The slip tendencies of both boreholes are summarised in a Mohr circle diagram in Figure 11. Fractures
 282 close to the failure line are critically stressed and a fluid injection shifts the Mohr circle and fracture
 283 planes, plotted as poles, closer to the failure line. The higher slip tendency of fractures oriented NW-

284 SE compared to N-S oriented fractures can be explained by their orientation relative to the input
 285 direction of the minimum horizontal stress.



286

287 *Figure 11. Visualization of the identified natural fractures in a Mohr circle. The fracture planes are*
 288 *colored according to the critical pore pressure increase needed to induce slip. The single fracture*
 289 *occurring in borehole SB2.2 is marked, the other fractures are from SB3.1. The solid line gives the Mohr-*
 290 *Coulomb criteria for fault reactivation. Dashed lines show the failure criterion after an increase in pore*
 291 *pressure through fluid injection, which reduces the effective stress and possibly reactivates optimally*
 292 *oriented faults. (Bröker, 2019)*



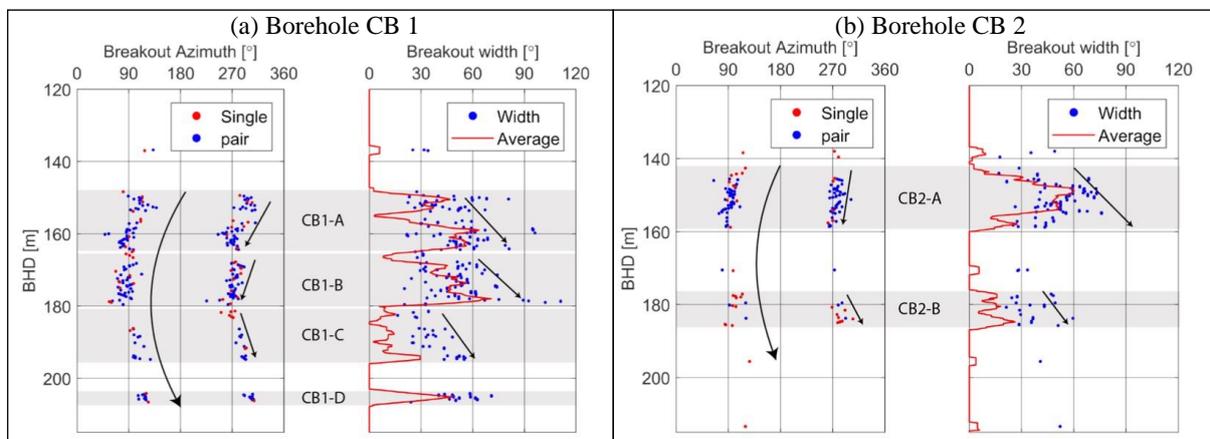
293

294 *Figure 12. Record from the dilatometer test at 9 m depth in SB2.2 (Bröker, 2019). Shown is the average*
 295 *deformation of the top, middle and bottom displacement transducer. The rising maximum pressure*
 296 *can be seen in all three loading cycles.*

297

298 The dilatometer test is used to obtain the modulus of deformation and modulus of elasticity of the in-
 299 situ rock mass. The tests were conducted at a depth of 9 m in Borehole SB2.2 and 22 m in Borehole
 300 SB4.1. The averaged deformation observed from the three transducers during the dilatometer test in
 301 SB2.2 can be seen in Figure 12. The first cycle pressurized the rock up to 3.4 MPa, the second one up to
 302 to 5.4 MPa and the final one up to 7.9 MPa. The average elastic modulus is, also considering the
 303 dilatometer test in the vertical borehole SB4.1, $42 \text{ MPa} \pm 18 \text{ MPa}$. The overall average of the
 304 deformation modulus from the loading cycle is $12 \text{ MPa} \pm 4 \text{ MPa}$ and from the reloading cycle 45 MPa
 305 $\pm 25 \text{ MPa}$. The high standard deviation shows the large scatter of the values.

306 ETHZ has conducted additional work in WP 2 on in situ stress analysis and borehole breakout analysis.
 307 The additional work is covered by part of the personnel hours of WP 4, due to the suspending of
 308 collaboration between ETH, GES, ESG in WP 4 (see explanation in WP 4). In these additional works, an
 309 azimuth log, tilt log, and caliper log were collected from the ATV logs of Boreholes CB1 and CB2.
 310 Breakouts along the boreholes CB1 and CB2 were identified based on the ATV logs: 262 breakouts
 311 were identified in CB1 between 120-220 m; 136 breakouts were identified in CB2 between 120-220 m
 312 (Figure 13). Discussions were made on variations in breakout presence, orientation and shape, and
 313 the underlying processes that cause these variations. Intersections between natural fractures and
 314 Boreholes CB 1 and CB 2 were identified using ATV logs and OTV logs. The direct link between borehole
 315 orientation, borehole tilt, and breakout presence is investigated, as well as the influence of the rate
 316 of change of the orientation and tilt. Estimations were made on the rock strength in areas with
 317 breakout presence.



318

319 *Figure 13. Breakout azimuth and width for Boreholes CB 1 and CB 2, zoomed into interval between*
 320 *120–220m BHD. Single breakouts and breakout pairs are indicated. Azimuth trend over the complete*
 321 *borehole is indicated, as well as trends in azimuth and width for groups of breakouts identified,*
 322 *CB1A–D and CB2A–B for CB1 and CB2 respectively. Breakout azimuth is given with respect to top of borehole*
 323 *(HS), orientation with respect to magnetic North (NM). (After Limborgh (2020))*

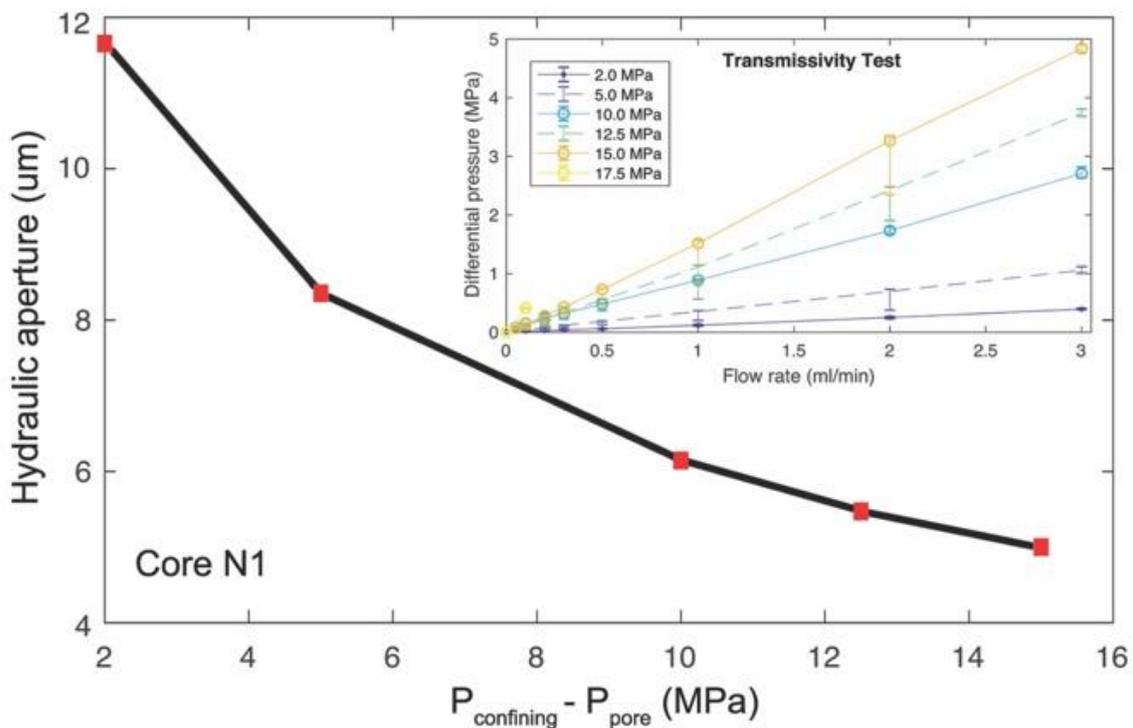
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325

326 3.2 Flow-through experiments and seismic risk analysis at BULGG

327 In WP 3: Stimulation and production optimization, ETH Zurich has conducted, fracture permeability
 328 characterization, mud cake formation, seismic risk analysis, and shear and acidizing stimulation.
 329 Similar to WP 2, due to the suspending of collaboration between ETH, GES, ESG in WP 4, the work of
 330 shear and acidizing stimulation used another part of the personnel hours. This deviation has also been
 331 approved by the Swiss Bundesamt für Energie (BFE).

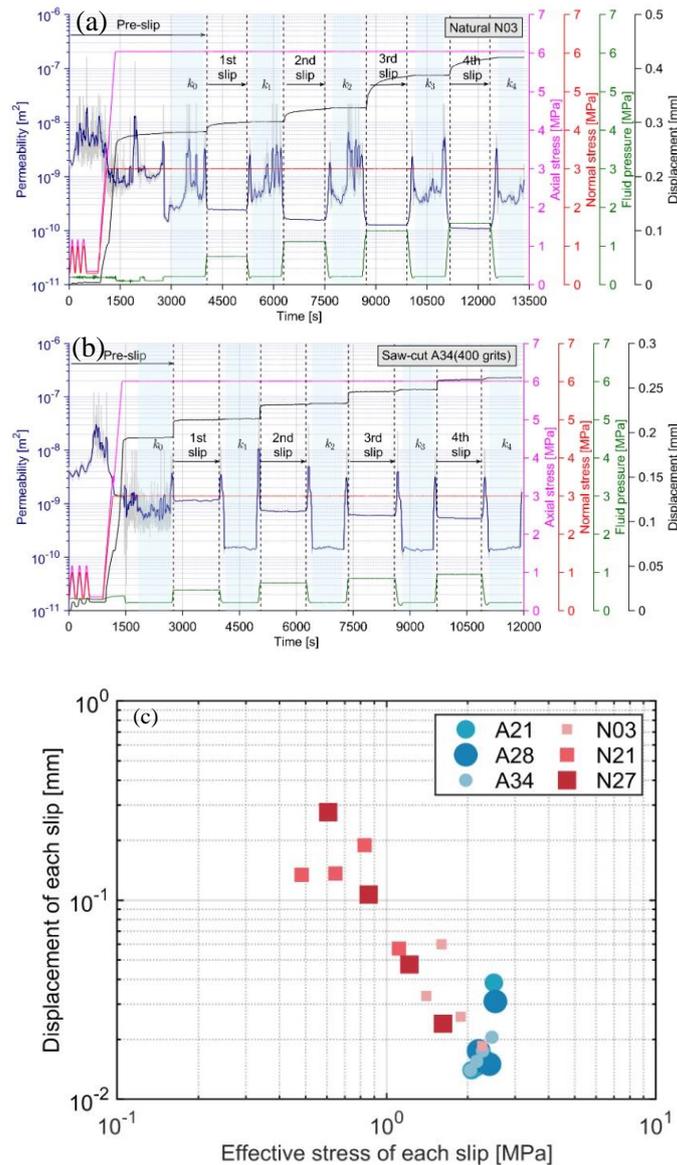
332 A series of the transmissivity tests has been conducted on naturally fractured Granodiorite specimens
 333 with a diameter of 2.8 cm and a length of 3.5 cm. The composition of the granodiorite specimens was
 334 determined, using Energy Dispersive X-Ray Analysis (EDX), to be: quartz, feldspar, biotite and chlorite.
 335 For each transmissivity measurement, a constant flow rate was set and the differential pressure (ΔP)
 336 between the inlet and the outlet of the specimen was recorded. Later, the confining pressure ($P_{confining}$)
 337 and the fluid pore pressure (P_{pore}) were varied to establish different aperture configurations. Finally,
 338 the evolution of the hydraulic aperture was mapped under different effective stresses ($P_{confining} - P_{pore}$)
 339 (Figure 14). These transmissivity results can be used to elucidate how fracture injectivity will be
 340 changed under different effective stresses.



341
 342 *Figure 14. Hydraulic aperture evolution under different effective stresses ($P_{confining} - P_{pore}$).*

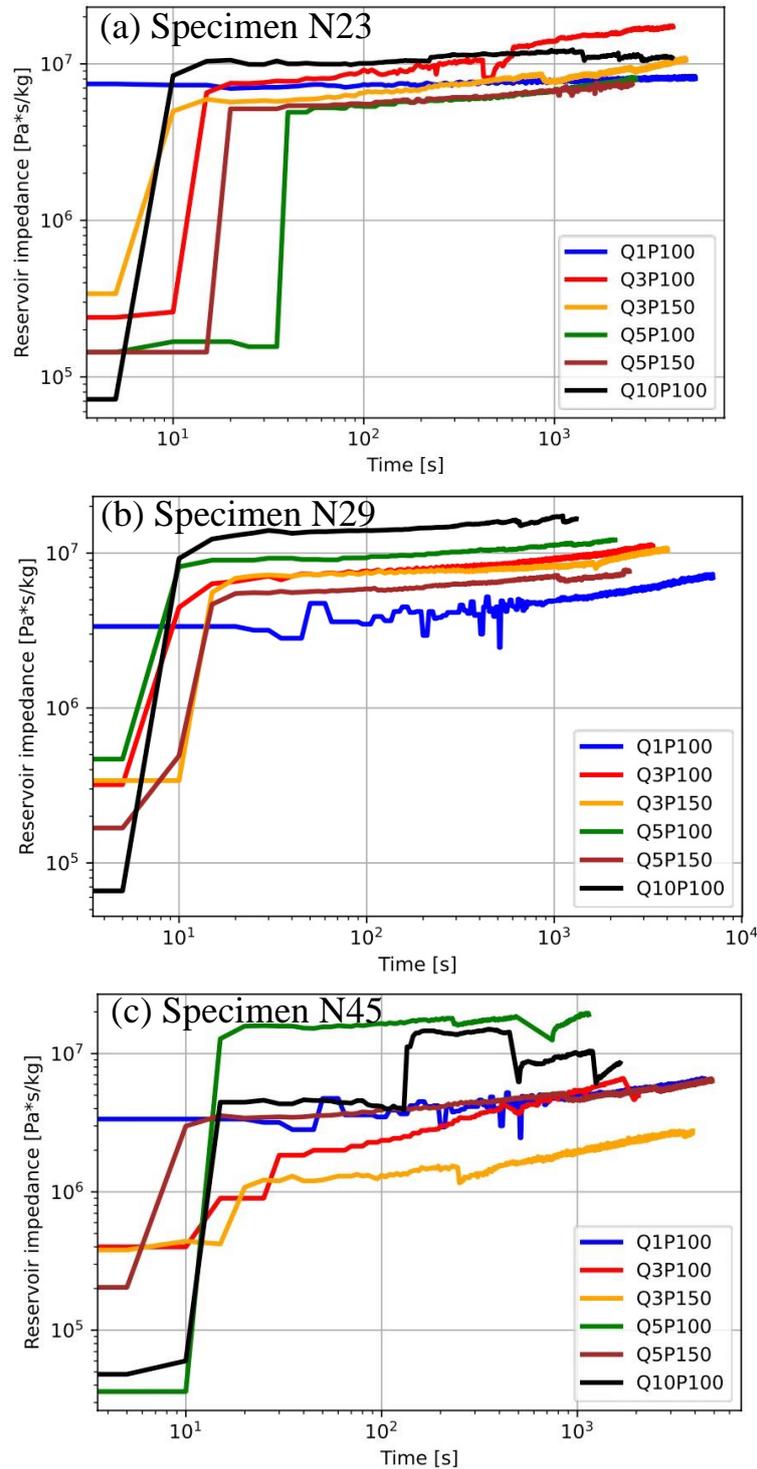
343 Fracture permeability is influenced by not only its effective stress, but also shear displacements. To
 344 examine how the permeability of fractured granite changes, when subjected to shearing events,
 345 induced by hydraulic stimulation, shear experiments were conducted to measure the permeability of
 346 granite samples with a single parallel natural/saw-cut fracture (Li et al., 2021). During the
 347 experiments, the shear slips were induced by increasing fluid pressure. For all examined specimens,
 348 the fracture permeability changes up to one order of magnitude after experiencing shear slips. No

349 consistent increasing or decreasing trend of permeability evolution was observed, neither for a single
 350 slip event, nor for the entire experiment. For most of the specimens, permeability increased during
 351 the first few shear slip events, reaching a maximum permeability after a critical shear displacement
 352 was achieved (0.0445 mm for Specimen N03, 0.0570 mm for Specimen N21, 0.0715 mm for Specimen
 353 N27, and 0.0310 mm for Specimen A28), exhibiting a positive relationship with respect to fracture
 354 roughness. Generally, permeability then decreased with further slip events after a large, in most cases
 355 the largest, shear displacement.



356

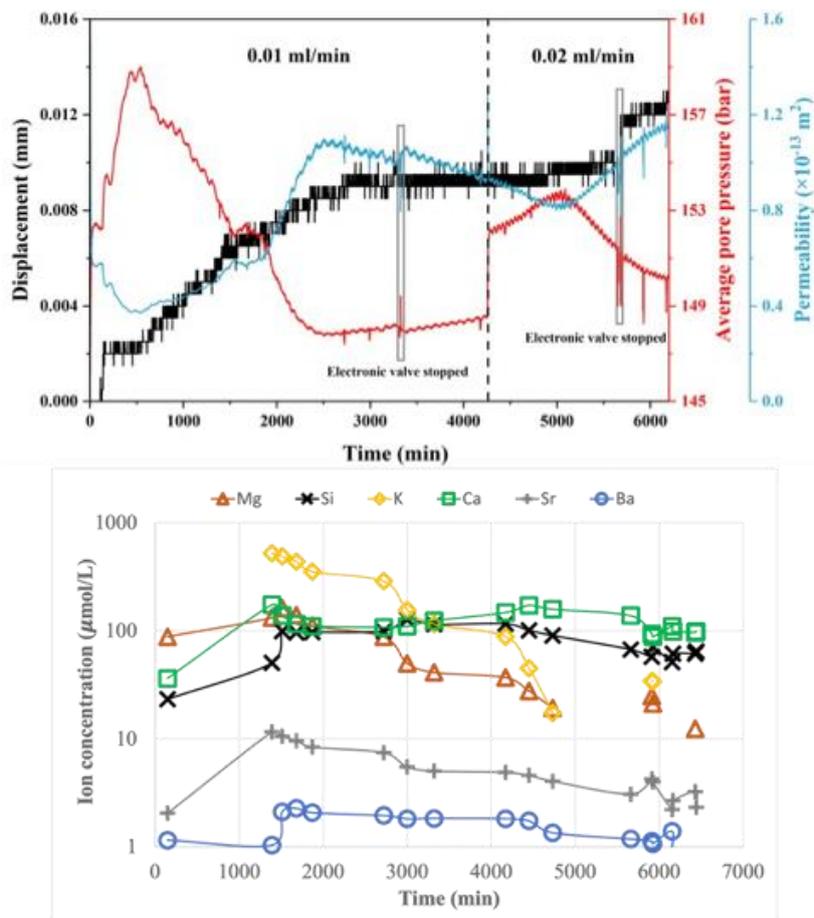
357 *Figure 15. Loading path, accumulated displacement (since the start of the experiment), and calculated*
 358 *permeability (original data in gray and filtered data in dark blue) of Specimens: (a) N03 and (b) A34.*
 359 *The range of shear displacements are marked by dotted lines. (c) Relationship between effective stress*
 360 *and shear displacement for each slip distance on a log-log plot. Specimens with larger fracture*
 361 *roughness values are marked by bigger markers. Specimens N03, N21 and N27 are natural fractures*
 362 *of different modes, and specimens A21, A28 and A34 are saw-cut fractures of different polishing*
 363 *finishes. (After Li et al. (2021))*



364

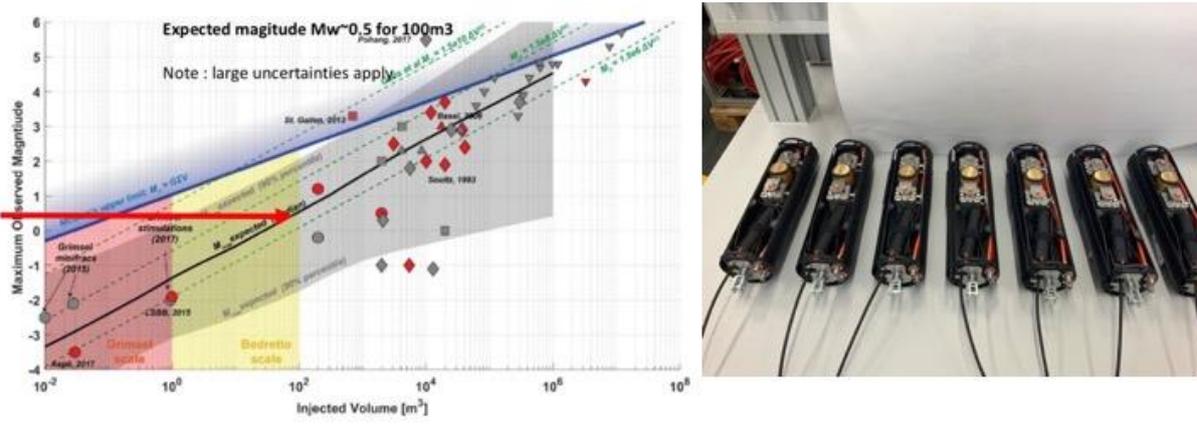
365 *Figure 16. Reservoir impedances for Specimens (a) N23, (b) N29, and (c) N45 during the formation of*
 366 *mud cake. In the legends, the numbers after Q indicate the flow rates with a unit of ml/min, and the*
 367 *numbers after P indicate the back pressures with a unit of bar during the experiment. (After Wang*
 368 *(2021))*

369 Mud cake often forms around wells when drilling mud infiltrates into the permeable zones. Mud cake
 370 provides a barrier to prevent further penetration and loss of drilling fluid, as well as a loss of produced
 371 fluids, into a permeable formation. It is thus necessary to quantitatively explore the associated
 372 phenomena when mud cake is formed within a fracture in a laboratory environment. ETH Zurich
 373 conducted a series of laboratory experiments focusing on mud cake formation on the fractures
 374 coupled with hydraulic and mechanical processes. Mud cake formation was monitored under reservoir
 375 conditions (2 kilometers). The potential effect of mud cake on the hydraulic properties of natural
 376 fractures was examined under pressure conditions (downstream pressure of 10 MPa, 15 MPa, and
 377 confining pressure of 40 MPa). Three specimens of naturally fractured granodiorite cores from the
 378 Grimsel Test Site in Switzerland were subjected to flow-through experiments at room temperature
 379 (23°C), due to the unexpected failure of the heating system. Fractures have been imaged using X-ray
 380 CT to characterize their surface roughness. Fracture transmissivities have been measured on various
 381 pore fluid pressures and effective stresses. Mud has been prepared with calcite (CaCO₃) particles of
 382 size around 60 micros. In total, 18 experiments have been conducted (Figure 16). Hydraulic aperture
 383 and reservoir impedance of fractures were evaluated during the formation of mud cake. In general,
 384 the higher injection rate, the higher the amount of mud being brought into the fractures, yielding
 385 larger reservoir impedance (Figure 16).



386
 387 *Figure 17. (Top) Shear displacements, average pore fluid pressure, and fracture permeability during*
 388 *the experiment. (Bottom) Major and tracer element concentrations during the experiment. Injection*
 389 *rate was kept at 0.01 ml/min till $t=4200$ min, and then increased to 0.02 ml/min for the rest of the*
 390 *experiment. (Kong et al., (2021))*

391 When fractures are critically stressed, acidizing stimulation could cause shear slip of fractures. ETH
 392 Zurich conducted additional work on shear and acidizing simulation on carbonate-filled fractures. A
 393 cylindrical specimen of 2.54 cm in diameter and 3.33 cm in length was sub-cored from a granite sample
 394 at a depth of 1897 m in Borehole EPS1 in Soultz, France. SEM analysis indicated that the filling minerals
 395 in the fracture are carbonates (mainly CaCO_3), quartz and hematite. The specimen was imaged before
 396 and after the shear and acidizing simulation using the customised X-ray CT at the GEG Group (Figure
 397 3). Before the shear and acidizing simulation, fracture permeability was measured to be $\sim 10^{-13} \text{ m}^2$ at a
 398 normal confinement of 28.5 MPa, a back pressure of 14.5 MPa, and a temperature of 100 ° C. Injection
 399 of 1.66 mol/L NaCl brine with 0.001 mol/L HCl was carried out for ~ 6500 min, under a normal
 400 confinement of 28.5 MPa, a back pressure of 14.5 MPa, and a temperature of 100 ° C. Shear
 401 displacement of the fracture was continuously monitored with a gauge meter: a total of 12
 402 micrometer displacement was observed during the 6500 min experiment (Figure 17). Effluents from
 403 the flow-thought experiment were collected and analyzed using ICPMS to obtain the ion
 404 concentrations for geochemical analysis (Kong et al., 2021). Gouge materials from the experiments
 405 have been collected and analysed. Using the monitored shear displacements, the combined rigidity of
 406 the fracture was estimated to be about 3.29 GPa, the seismic moment magnitude was calculated as
 407 6.9 Nm, the standard seismic moment was determined as 60.8 Nm, and the maximum seismic
 408 moments of the two continuous slip events (during injection rate of 0.01 ml/min and 0.02 ml/min) can
 409 be computed as 18.61 Nm and 4.96 Nm, respectively (Kong et al., 2021).



410
 411 *Figure 18. (Left) Seismicity estimation in Grimsel scale and Bedretto scale (Ma, 2019). (Right): Seismic*
 412 *monitoring network (Streamer 2.0) for seismic activity monitoring.*

413 Seismicity is often associated with fluid injection at depth, including disposal of wastewater into deep
 414 formations, hydraulic fracturing of shale formations, and reservoir creation during the development
 415 of enhanced geothermal systems. It has been reported that the maximum seismic moment, induced
 416 by fluid injection, is limited to the volume of injected fluid times the modulus of the rock rigidity.
 417 According to this principle, given the expected injection fluid volume of a maximum of 100 m³ in the
 418 BUL, we estimate a maximum event of 0.5 Mw (Figure 18). Note that large uncertainties apply to this
 419 estimation, due to the uncertainty of rock heterogeneity. To eliminate the theoretical uncertainties
 420 and support long-term seismic risk assessment in the BULGG, seismic monitoring is proposed for the
 421 BULGG. This monitoring network covers micro-seismic monitoring at scales of a hundred meters to
 422 about 10 km distance from the site, with typically earthquake magnitudes of 0.5 to 3.0 Mw. ETHZ also
 423 contributed to the design and construction of the sensors (Streamer 2.0) of the seismic monitoring
 424 network (Figure 18).

425 3.3 Collaboration between ETHZ and ESG

426 In WP 4: Environmental protection, ETHZ was planning to collaborate with GES and ESG on the
427 effluents collection and measurements of the Rittershoffen site. A meeting between ETHZ (Dr.
428 Xiangzhao Kong) and ESG (Dr. Nicolas Cuenot) took place at ETHZ on January 16, 2020. During that
429 meeting, a general agreement has been made that ESG and ETHZ agreed to work together to carry out
430 the planned tasks. On April 8, 2020, a detailed collaboration plan for Tasks 4.2 and 4.3 was proposed
431 by ETHZ (Dr. Xiang-Zhao Kong) between ETHZ (Dr. Xiang-Zhao Kong) and ESG (Dr. Nicolas Cuenot):

- 432 (1) geochemical equilibrium calculations on the measured gas concentrations and species
433 concentrations;
- 434 (2) investigation of the gas composition discrepancy between measurements of the effluents
435 and the emitted gases at the venting valves;
- 436 (3) geochemical equilibrium calculations (mainly super-saturation of potential minerals) on the
437 deposit scales at the heat exchangers;
- 438 (4) geochemical equilibrium calculations on scale inhibitors to control scaling.

439 To carry out this collaboration plan, ESG agreed to share with ETHZ the measurement data from the
440 Rittershoffen site. ETHZ thus started preparing the data transfer agreement. However, from that time
441 on, there has been no further response from ESG (possibly because of the COVID-19 pandemic),
442 despite several (~5) contact attempts over several months following April 8, 2020, by Drs. Kong and
443 Saar to contact ESG. In June 2021, ETH learned that Dr. Nicolas Cuenot was leaving ESG. Despite that
444 ETHZ wanted to conduct the proposed tasks with ESG at the Rittershoffen site, the collaboration in
445 WP 4 on Tasks 4.2 and 4.3 between ESG and ETHZ was unfortunately not possible to be carried out,
446 given the unfortunate and difficult circumstances. In total, ETHZ has spent 49.2 personnel hours on
447 WP4. The rest of the personnel hours of WP4 (about 944 hours) has been used to carry out the
448 additional works in WP 2 and WP 3, as previously explained in paragraphs above.

449 4. Conclusions

450 ETH Zurich has worked on fracture analysis, stress analysis, injectivity test, measurements of
451 deformation modulus and elasticity modulus, and borehole breakout analysis for several boreholes at
452 BULGG, using various methods, including Acoustic Televiewer (ATV) and Optical Televiewer (OTV)
453 investigation, mini-frac tests and dry (packer) reopening tests, step-rate injection tests, dilatometer
454 tests, sleeve fracturing method, fracture compliance method, G-function, square root of time, bilinear
455 pressure-decay, and jacking pressure analyses techniques. Various important parameters have been
456 obtained, including BULGG in-situ stress, (rock matrix and fracture) injectivity, deformation modulus,
457 elasticity modulus, pore fluid pressure, slip tendency, fracture density, fracture orientation, and etc.
458 Various flow-through experiments of ETHZ yielded critical insights in the relationship between fracture
459 transmissivity and the applied mechanical stress, thermal stress, and chemical stimulation. In
460 particular, ETHZ has performed experiments of mud infiltration into fractures to examine the mud
461 cake formation process. The shear-slip experiments yielded critical estimates on fracture combined
462 rigidity, seismic moment magnitude, standard seismic moment, and maximum seismic moment. ETHZ
463 also gave estimates on fluid injection-induced seismicity at BULGG.

464

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